## <sup>1</sup> Wet scavenging process of particulate matter $(PM_{10})$ : A <sup>2</sup> multivariate complex network approach

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#### 12 Abstract

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This paper reports the results of research on  $PM_{10}$  wet scavenging by 13 rainfall using a new multilayer complex networks called Multiplex Visibility 14 Graphs (MVG). To the best of our knowledge, this work is the first to assess 15  $PM_{10}$  wet deposition using multivariate time series according to African 16 dust seasonality. We considered 11 years of daily  $PM_{10}$  and rainfall data 17 from the Guadeloupe archipelago. To analyse the impact of rainfall on  $PM_{10}$ 18 behaviour, two MVG parameters were computed: the average edge overlap 19  $(\omega)$  and the interlayer mutual information  $(I_{PM_{10}Rainfall})$ . On the 1-d scale, 20 the  $\omega$  results showed that the wet scavenging process was higher during the 21 second half of the year when the high dust season and the rainy season are 22 juxtaposed. This highlights a greater correlation between the microscopic 23 structure of the signal, and the impact of rainfall on  $PM_{10}$  concentrations 24 is more significant when the atmosphere is loaded with dust. The joint 25 probability computed between the  $PM_{10}$  and rainfall nodes confirmed this 26 trend. The  $I_{PM_{10}Rainfall}$  results indicated a correlation between  $PM_{10}$  and 27

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rainfall structures throughout the year. Furthermore,  $I_{PM_{10}Rainfall}$  values were higher during the transition periods between winter and summer (and vice versa). Our study showed that MVG is a powerful technique for investigating the relationship between at least two nonlinear time series using a multivariate time series.

<sup>33</sup> Keywords:  $PM_{10}$ , Wet scavenging, Multiplex visibility graphs, Complex <sup>34</sup> networks, Caribbean area

#### 35 1. Introduction

In geoscience, precipitation is a key component of the water cycle (Schnei-36 der et al., 2014) and of atmospheric circulation (Kidd and Huffman, 2011). 37 In recent decades, the removal of atmospheric Particulate Matter (PM) by 38 falling precipitation has greatly interested the scientific community (González 39 and Aristizábal, 2012; Ouyang et al., 2015; Wu et al., 2018). This phe-40 nomenon, which can occur through liquid (rain) and solid (snow) forms of 41 precipitation, is called "wet deposition" (Kim et al., 2012; Singh et al., 2016). 42 Numerous studies have shown that wet scavenging of PM by rainfall is one 43 of the primary precipitation processes for wet deposition (Laouali et al., 44 2012; Tiwari et al., 2012; Yoo et al., 2014; Singh et al., 2016; Olszowski, 45 2017; Wu et al., 2018; McClintock et al., 2019). Raindrops falling through 46 the air column, bump into and collect air particles. Raindrops approach the 47 particles, apply a force via the air as a medium, and change trajectory (Son-48 wani and Kulshrestha, 2019). The collision between the raindrops and the 49 PM is conditioned by size and relative location (Olszowski, 2017). The two 50 primary wet scavenging mechanisms related to rainfall are rainout (in-cloud 51 scavenging) and washout (below-cloud scavenging) (Dallarosa et al., 2005; 52 Tombette et al., 2009; Sonwani and Kulshrestha, 2019). Studies have shown 53

that wet scavenging can remove 30% of the aerosols from the troposphere
(Murakami et al., 1983; Schumann, 1989).

Over the past decades, two types of PM have received special attention 56 due to their health impact: fine particles (particulate diameter  $< 2.5 \ \mu m$ , 57  $PM_{2.5}$ ) and coarse particles (particles with diameters between 2.5 and 10 58  $\mu$ m,  $PM_{10-2.5}$ ) (Bayraktar et al., 2010; Plocoste and Calif, 2019). Epidemi-59 ological studies reveal that short- and long-term exposure to high concen-60 trations of  $PM_{2.5}$  and  $PM_{10}$  can cause human health problems (Weinmayr 61 et al., 2010; Atkinson et al., 2014; Lu et al., 2015). The authors focused on 62  $PM_{10}$ , which also strongly impacts climate (Plocoste and Pavón-Domínguez, 63 2020b; Plocoste et al., 2020a). 64

In the Caribbean, air quality is frequently degraded by African dust 65 (Euphrasie-Clotilde et al., 2020). Dust haze episodes primarily occur dur-66 ing summer (Petit et al., 2005; Prospero et al., 2014). Many studies present 67 the processes that allow the transport of dust over the Atlantic Ocean (Perry 68 et al., 1997; Prospero, 1999; Prospero and Lamb, 2003; Engelstaedter et al., 69 2006; Kumar et al., 2014; Euphrasie-Clotilde et al., 2020). As the wet scav-70 enging of  $PM_{10}$  is a standard indicator of air quality in a given area, the 71 aim of this study was to investigate the wet scavenging process of  $PM_{10}$  by 72 rainfall in the Caribbean Basin. Additionally, we aim to determine whether 73 there is a link between wet scavenging efficiency and African dust seasonal-74 75 ity.

A newly developed method termed Multiplex Visibility Graph (MVG) was used to perform this study (Lacasa et al., 2015). The methodology is based on a previous technique called Visibility Graph (VG), first introduced by Lacasa et al. (2008). The main idea is to transform a time series into a complex network, which can later be analysed, and to preserve some of the

original information. Most of the variants of this method focus on analysis 81 of a single time series (Luque et al., 2009; Lan et al., 2015; Carmona-Cabezas 82 et al., 2019b; Iacovacci and Lacasa, 2019) and have been applied to applica-83 tions related to univariate time series (Mali et al., 2018; Carmona-Cabezas 84 et al., 2019a; Plocoste et al., 2021b). However, owing to their stochastic 85 properties, atmospheric processes are frequently related to numerous de-86 grees of freedom; that is, their behaviour is governed by a multivariate time 87 series. To overcome this drawback, the MVG technique applies the visibility 88 approach to examine nonlinear multivariate time series (Lacasa et al., 2015). 80 After transforming the time series into complex networks, the results were 90 used to build a multi-layered structure that could be analysed. Owing to 91 recent advances in the theory of multilayer networks (Bianconi, 2013; Kivelä 92 et al., 2014; Battiston et al., 2014; Lacasa et al., 2015), additional informa-93 tion can be retrieved from the original multivariate time series. To the best 94 of our knowledge, no study has yet investigated  $PM_{10}$  wet scavenging using 95 a multivariate time series. Here, 11 years of daily  $PM_{10}$  and rainfall data 96 from the Guadeloupe archipelago were analysed. 97

#### 98 2. Site and data collection

<sup>99</sup> The Guadeloupe archipelago  $(16.25^{\circ}N - 61.58^{\circ}W)$  is a French overseas <sup>100</sup> region located in the central Caribbean Basin (Plocoste et al., 2019). The <sup>101</sup> small territory (~1,800 km<sup>2</sup>; 390,250 inhabitants) has an insular tropical <sup>102</sup> climate with meteorological characteristics that vary by location due to mi-<sup>103</sup> croclimates (Bertin and Frangi, 2013). According to the Köppen-Geiger <sup>104</sup> climate classification (Peel et al., 2007), Guadeloupe is in the "Af (tropical <sup>105</sup> rainforest)" category.

For this study, time series of Particulate Matter  $(PM_{10})$  and rainfall were 106 used. Hourly  $PM_{10}$  data were provided by Gwad'Air Agency (http://www.gwadair.fr/), 107 which manages the Guadeloupe air quality network.  $PM_{10}$  concentrations 108 were measured using the Thermo Scientific Tapered Element Oscillating Mi-109 crobalance (TEOM) models 1400ab and 1400-FDMS. From 2005 to 2017, 110 the air quality network-principally located at the centre of the island-has 111 only one  $PM_{10}$  sensor at Pointe-à-Pitre (16.2422°N 61.5414°W) from 2005 112 to 2012 and at Baie-Mahault (16.2561°N 61.5903°W) since 2015. Because of 113 the proximity between the air quality stations ( $\sim 5.5 \text{ km}$ ),  $PM_{10}$  measure-114 ments were performed under the same environmental conditions. Rainfall 115 measurements were made by Météo France at the international airport of 116 Pôle Caraïbes at Abymes (16.2630°N 61.5147°W) using a Precis-Mecanique 117 3070. As with the  $PM_{10}$  time series, the Météo France observations are an 118 hourly rainfall time series. Both measurements were made in the insular 119 continental regime (Plocoste et al., 2018; Plocoste and Pavón-Domínguez, 120 2020a). To assess the possible wet scavenging phenomenon over an entire 121 day, hourly  $PM_{10}$  data were converted into daily average values, whereas 122 rainfall data were converted into daily average and daily sum values. By 123 computing the Pearson correlation coefficient between the daily average 124  $PM_{10}$  and the daily rainfall data (sum then average), the same result was 125 obtained (R = -0.14). Many studies demonstrate the cumulative effect of 126 rainfall on atmospheric processes (Winstanley, 1973; Johnson and Ciesiel-127 ski, 2000). In addition, to account for hours with and without rainfall (0 128 mm) in a day, the authors favoured the sum over the average for the stochas-129 tic analysis. Thus, 11 years of simultaneous measurements between the daily 130 average  $PM_{10}$  and the daily sum of rainfall were available for this study (a 131 total of 3,849 points per time series). Figure 1 shows the sequence of the 132

analysed time series. A slight lag appears to exist between the groups ofpeaks.



Figure 1: Illustration of simultaneous measurement sequences between (a) daily average  $PM_{10}$  concentrations and (b) the daily sum of rainfall between 2005 and 2011.

#### 135 3. Theoretical framework

#### 136 3.1. Visibility graphs

A graph is a mathematical object composed of a set of vertices (or nodes) which are connected by a set of lines or edges. A relatively recent tool called Visibility Graph (VG) allows the transformation of two-dimensional

sets of points into graphs or networks (Lacasa et al., 2008). VG has great 140 applicability for time series analysis and produces networks that inherit 141 many of the properties of the original time series (Lacasa and Toral, 2010). 142 The points in the time series correspond to the nodes in the graph. The 143 edges of the graph (which connect nodes) are selected by checking which 144 pairs of points meet the visibility criterion, which is as follows: two points 145 from the time series  $(t_a, y_a)$  and  $(t_b, y_b)$  are connected only if any other point 146  $(t_c, y_c)$  located between them  $(t_a < t_c < t_b)$  fulfils the following relationship 147 (Lacasa et al., 2008): 148

$$y_c < y_a + (y_b - y_a) \frac{t_c - t_a}{t_b - t_a}$$
(1)

To construct the graph, this algorithm was applied to every pair of points in the signal. Two consecutive nodes are always connected, because there are no intermediate points.

A graph is commonly expressed via its adjacency matrix, whose rows 152 store the information of each node. If an element  $a_{ij}$  is equal to 1, nodes i 153 and j are connected, the opposite is true if  $a_{ij}$  is equal to 0. In the case of a 154 time series with N points, the resulting VG is represented by an N×N adja-155 cency matrix, which has special properties that facilitate computation; the 156 adjacency matrix is symmetric  $(a_{ij} = a_{ji})$  and hollow  $(a_{ii} = 0)$ ; additionally, 157 all the nearest neighbours are visible to one other  $(a_{ij} = 1 \text{ for } j = i \pm 1)$ . 158 In general, the adjacency matrix has the following form (Carmona-Cabezas 159

160 et al., 2019a):

$$V = \begin{pmatrix} 0 & 1 & \cdots & a_{1,N} \\ 1 & 0 & 1 & \vdots \\ \vdots & 1 & \ddots & 1 \\ a_{N,1} & \cdots & 1 & 0 \end{pmatrix}$$
(2)

#### 161 3.2. Degree centrality

Degree is the most commonly used of the principal properties that can be studied from a graph and one of the centrality parameters used to measure the importance of different nodes in the graph with relation to the rest of them, using different criteria (Latora et al., 2017). The degree of a node  $(k_i)$ measures (in an undirected graph) the number of nodes that are reciprocally connected to a given node. By considering the adjacency matrix, the degree can be computed as  $k_i = \sum_j a_{ij}$ .

Once the degree of every point is computed, a degree probability dis-169 tribution P(k) can be obtained for the graph (here, VG). P(k) accounts 170 for the probability of having each value of degree in the graph. To obtain 171 information on the nature of the series, the degree distribution is analysed 172 (Lacasa et al., 2008; Mali et al., 2018; Pierini et al., 2012). If the right tail of 173 the degree distribution (for high values of degree) can be fitted by a power 174 law such as  $P(k) \propto k^{-\gamma}$ , the time series has a fractal nature (Lacasa et al., 175 2008). This part of the distribution is related to the hubs (nodes with the 176 highest degrees), which are, by definition, rare in a graph. The exponent in 177 the power law is the coefficient, which is related to the Hurst exponent in 178 series related to Brownian motion (Lacasa et al., 2009). 179

#### 180 3.3. Multiplex visibility graph

Another application of VGs in the context of multivariate analysis, is 181 the use of multi-layered networks. This methodology was recently intro-182 duced as Multiplex Visibility Graph (MVG) (Lacasa et al., 2015). The 183 main idea behind MVG is to build each of the layers M with VGs from 184 the different variables of the study. Therefore, as VG is represented by its 185 adjacency matrix, so MVG is identified by a vector of adjacency matrices 186  $\Omega = \{A^{[1]}, A^{[2]}, ..., A^{[M]}\}.$  In the last expression,  $A^{[\alpha]}$  corresponds to the VG 187 adjacency matrix of the VG in the  $\alpha$ -dimension (or layer in the multiplex), 188 which comes from the  $\alpha$  variable of the multivariate time series (see Figure 2, 189 where  $PM_{10}$  and rainfall sample time series are transformed for illustrative 190 purposes). 191



Figure 2:  $PM_{10}$  and rainfall time series (left) are converted into complex networks with the VG algorithm (centre), which is described by an adjacency matrix  $(A^{PM_{10}} \text{ and } A^{Rainfall})$ . Then, both are combined to design a two-layered MVG, called  $\Omega$  (right image).

After construction, MVG is analysed to obtain information regarding the 192 system of the time series. The two measures used for such purposes (Nicosia 193 and Latora, 2015) and chosen for this work are Average Edge Overlap ( $\omega$ ) 194 and Interlayer Mutual Information  $(I_{\alpha,\beta})$ .  $\omega$  averages the number of layers 195 on which a given edge between a pair of nodes can be found.  $I_{\alpha,\beta}$  measure 196 the correlations between the degree distributions of the given layers  $\alpha$  and 197  $\beta$ . In this study, the layers correspond to daily  $PM_{10}$  concentrations and 198 total rainfall. 199

The computation is relatively straightforward after the MVG and degree distributions of each layer are obtained. Equation 3 shows the formula to compute the  $\omega$  of a given MVG (Lacasa et al., 2015):

$$\omega = \frac{\sum_{i} \sum_{j>i} \sum_{\alpha} a_{ij}^{[\alpha]}}{M \sum_{i} \sum_{j>i} \left(1 - \delta_{0, \sum_{\alpha} a_{ij}^{[\alpha]}}\right)}$$
(3)

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All quantities were previously defined in the text;  $\delta_{0,\sum_{\alpha}a_{ij}^{[\alpha]}}$  corresponds to 204 a Kronecker delta, which is 1 when  $\sum_{\alpha} a_{ij}^{[\alpha]}$  is null, and otherwise 0. The 205 maximum value of  $\omega = 1$  indicates that all the layers and, therefore, the time 206 series are identical. Conversely, the minimum possible value of  $\omega = 1/M$ 207 indicates that every edge in the MVG can be found only in a singular layer. 208 Overall, this quantity provides an idea of the expected number of layers on 209 which an edge can be found. In addition, a high  $\omega$  value indicates a high 210 correlation in the microscopic structure of the signal (Lacasa et al., 2015). 211 212 Additionally,  $I_{\alpha,\beta}$  is defined in Equation 4 (Lacasa et al., 2015):

$$I_{\alpha,\beta} = \sum_{k^{[\alpha]}} \sum_{k^{[\beta]}} P(k^{[\alpha]}, k^{[\beta]}) \log \frac{P(k^{[\alpha]}, k^{[\beta]})}{P(k^{[\alpha]})P(k^{[\beta]})}$$
(4)

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where  $P(k^{[\alpha]}, k^{[\beta]})$  is the joint probability of having a degree of  $k^{[\alpha]}$  in layers  $\alpha$  and  $k^{[\beta]}$  in layer  $\beta$  that can be obtained using the following formula:

$$P(k^{[\alpha]}, k^{[\beta]}) = \frac{N_{k^{[\alpha]}, k^{[\beta]}}}{N}$$
(5)

where  $N_{k^{[\alpha]},k^{[\beta]}}$  is the number of nodes with a degree of  $k^{[\alpha]}$  in layer  $\alpha$ and  $k^{[\beta]}$  in layer  $\beta$ ;  $N_{k^{[\alpha]},k^{[\beta]}}$  is divided by N, which is the total number of nodes or points in the time series.

#### 219 4. Results and Discussion

#### 220 4.1. Preliminary analysis

We conducted a preliminary analysis to investigate  $PM_{10}$  and rainfall 221 seasonality throughout the year. Figure 3 illustrates the monthly average 222  $PM_{10}$  concentrations and the monthly summation of rainfall data over 11-y 223 period. Seasonality is observed in both curves, with a high dust season from 224 May to September (Plocoste and Pavón-Domínguez, 2020b) and a rainy 225 season from July to November (Bertin and Frangi, 2013). Van der Does 226 et al. (2020) observed the same behaviour for both parameters in Barbados. 227 The Inter-Tropical Convergence Zone (ITCZ) dynamics throughout the year 228 play a key role in these seasonal behaviours. In summer, the activation of 229 dust sources from the Saharan and Sahelian deserts coupled with the up-230 ward northward movement of the ITCZ  $(10 - 20^{\circ}N)$  (Moulin et al., 1997; 231 Adams et al., 2012; Euphrasie-Clotilde et al., 2020) allows the transport of 232 dust plumes from the African coast to the Caribbean area (Petit et al., 2005; 233 Prospero et al., 2014; Euphrasie-Clotilde et al., 2021). According to a statis-234 tical study lasting over a decade (Plocoste et al., 2020b), average  $PM_{10}$  and 235 kurtosis are 1.5 times higher and 5.5 times lower during the high dust season, 236

respectively, due to the recurrence of dust plumes compared with the low 237 dust season. From October to April,  $PM_{10}$  concentrations are primarily re-238 lated to marine aerosols (Clergue et al., 2015; Rastelli et al., 2017), because 239 of the insular context of the Guadeloupe archipelago. These aerosols are 240 advected by trade winds which blow continuously from east to west across 241 the Atlantic Ocean (Plocoste et al., 2014; Plocoste and Pavón-Domínguez, 242 2020a). Consequently, the contribution of marine aerosols to  $PM_{10}$  con-243 centrations remains constant throughout the year. Figure 3 shows that the 244 standard deviations exhibit their lowest values from October to April. Thus, 245 marine aerosols are one of the primary constituents of the  $PM_{10}$  background 246 atmosphere (Plocoste et al., 2021a). The ITCZ movement toward the north 247 generates precipitation carried by trade winds during the boreal summer 248 (the rainy season) (Giannini et al., 2000; Muñoz et al., 2008). During the 249 boreal winter (mid-January to March), the ITCZ awakens the Azores anti-250 cyclone due to its southerly movement, which reduces cloud generation (the 251 dry season) (Bertin and Frangi, 2013). 252

#### 253 4.2. Degree distribution

#### 254 4.2.1. Overall analysis

Before performing a profound analysis of the impact of rainfall on  $PM_{10}$ 255 concentrations in the MVG frame, both time series were analysed sepa-256 rately in the VG frame. The classical first approach is to study the degree 257 distribution P(k) of each time series. Figure 4(a) and 4(b) show the degree 258 distributions obtained for the  $PM_{10}$  and rainfall time series over the 11-y 259 period, respectively. Both plots highlight the fractal nature of the time se-260 ries. The tail region of P(k) in the log-log plot can be fitted by a power 261 law, such as  $P(k) \propto k^{-\gamma}$ , where  $\gamma_{PM_{10}}$  and  $\gamma_{rainfall}$  equal 3.11, and 2.84, 262



Figure 3: Monthly conditional average of  $PM_{10}$  and rainfall time series over the 11-y period; here, a representative year for one decade. The whiskers depict the standard deviations.

respectively. In the literature, the fractal nature of the  $PM_{10}$  (Dong et al., 264 2017; Nikolopoulos et al., 2019; Plocoste et al., 2021b) and rainfall (Olsson 265 et al., 1993; Breslin and Belward, 1999; Maskey et al., 2016) data has been 266 observed.

To assess the behaviour of the highest degree (so-called hubs), the time 267 series values versus their degrees (v-k plot) were analysed for each parameter 268 as introduced by Pierini et al. (2012). Figure 4(c) and 4(d) illustrate the v-k 269 plot for the  $PM_{10}$  and rainfall data, respectively. In both cases, the hubs 270 were related to the highest values of each time series. Carmona-Cabezas 271 et al. (2019a) found the same tendency for hubs of a tropospheric ozone 272 time series in Cadiz, Spain. The value of precipitation has an almost linear 273 relationship to the degree of rainfall. Thus, the degree of the rainfall nodes 274 can be used to identify both high and low rainfall values. In addition, the 275



Figure 4: At the top, degree distribution of the visibility graph for (a)  $PM_{10}$  and (b) rainfall in a log-log plot for all the data. A the bottom, the relationship between the time series values and their degrees in (c)  $PM_{10}$  and (d) rainfall.

276  $PM_{10}$  dot distribution appears more heterogeneous, because of the wide 277 annual variability of African dust haze (Plocoste et al., 2017, 2020a).

#### 278 4.2.2. Monthly analysis

We use the first centrality measure (degree centrality) to study the importance of the node for  $PM_{10}$  and rainfall throughout the year (Carmona-Cabezas et al., 2019a). Figure 5(a) and 5(b) highlight the monthly behaviour of the average degree and standard deviation from the degree distribution of

the  $PM_{10}$  and rainfall time series. A trend merged in both curves. The decay 283 of  $PM_{10}$  hubs begins at the onset of the high dust season (May-September) 284 (Plocoste and Pavón-Domínguez, 2020b; Plocoste et al., 2021b); the decay 285 for rainfall hubs begins at the onset of the hurricane season (June–October) 286 (Tartaglione et al., 2003; Dunion, 2011) in the Caribbean Basin. Figure 3 287 shows that the monthly behaviour of  $PM_{10}$  and rainfall over the period of a 288 decade confirms this trend, increases in  $PM_{10}$  and rainfall begin in May and 289 June, respectively. The above-mentioned results show the impact of season-290 ality on node distribution and highlight that the VG frame is sensitive to 291 time-series behaviour. 292

#### 293 4.3. Multiplex visibility graph

After performing the analysis of the  $PM_{10}$  and rainfall univariate time 294 series in the VG frame (i.e., transformation of the time series into a complex 295 network), both complex networks were combined to design a two-layered 296 multivariate network. Here, we investigate the wet scavenging process of 297  $PM_{10}$  by rainfall. The authors focused on two approaches, which demon-298 strate the abundance of single edges across layers (average edge overlap) 299 and the presence of interlayer correlations of the node degrees (interlayer 300 mutual information) (Lacasa et al., 2015; Nicosia and Latora, 2015). The 301 first approach measures the overall coherence in the multivariate time series, 302 and the second evaluates structural correlation. 303

#### 304 4.3.1. Average edge overlap analysis

Figure 6 illustrates the monthly average edge overlap values ( $\omega$ ) and their standard deviations over the 11-y period.  $\omega > 1/M$  and  $\omega < 1$ ; thus, the two layers are different, and edges can be found in both layers. These



Figure 5: Computed average degree and standard deviation from the degree distribution of each month over the 11-y period for (a)  $PM_{10}$  and (b) rainfall. Each monthly value is the average of the computed 11-y values.

two criteria prove that there is an interaction between  $PM_{10}$  and rainfall in the MVG frame.  $\omega$  is a sensitive parameter with small variations (increases and decreases) (Lacasa et al., 2015). Here,  $\omega$  was almost constant from January to June.  $\omega$  values are higher from July to December and peaked in September, indicating a greater correlation in the microscopic structure of the signal (Lacasa et al., 2015; Carmona-Cabezas et al., 2020) that repre-

sents interaction on the order of 1-d, which is the minimum time resolution. 314 Therefore, at the 1-d scale, the wet scavenging process of  $PM_{10}$  by rainfall is 315 more significant during the last six months of the year. A 20-y precipitation 316 study in the Luquillo Mountains of Puerto Rico McClintock et al. (2019) 317 also found a summer maximum in wet dust deposition. Physically, these 318 findings make sense, as the summer corresponds to the high dust (Plocoste 319 and Pavón-Domínguez, 2020b) and rainy (Bertin and Frangi, 2013) seasons 320 in the Caribbean Basin. The atmosphere is loaded with dust, and the im-321 pact of rainfall on  $PM_{10}$  is greater. Tiwari et al. (2012) observed that low 322  $PM_{10}$  concentrations in New Delhi occurred during the monsoon (August-323 September) season due to the washout phenomenon. A 10-y study of air 324 pollutants  $(PM_{10}, CO, NO_2, SO_2, \text{ and } O_3)$  and precipitation over South 325 Korea highlighted that  $PM_{10}$  is most effectively scavenged by summertime 326 rainfall due to its particulate nature (Yoo et al., 2014). Because of data 327 availability, determining which wet scavenging process (rainout or washout) 328 is more efficient is difficult (Pillai et al., 2002; Tombette et al., 2009; Bayrak-329 tar et al., 2010). According to Sonwani and Kulshrestha (2019), the level 330 of aerosols in and under clouds at the time of precipitation is crucial, as it 331 determines whether both phenomena occur simultaneously. 332

During the high dust season, the wet scavenging phenomenon naturally 333 reduces  $PM_{10}$  concentrations in the atmosphere. Due to the impact on res-334 piratory and cardiovascular diseases, diminishing  $PM_{10}$  concentrations after 335 dust outbreaks is crucial (Gurung et al., 2017; Zhang et al., 2017; Momtazan 336 et al., 2019; Feng et al., 2019). African Easterly Waves (AEWs) (Prospero 337 and Carlson, 1981; Plocoste et al., 2021a), which precede and follow the 338 dust plumes, are the principal generator of precipitation during the high 339 dust period (Dominguez et al., 2020) and regulate  $PM_{10}$  concentrations in 340

341 the Caribbean area.



Figure 6: 11-y monthly average edge overlap values ( $\omega$ ) and their standard deviations.

#### 342 4.3.2. Interlayer mutual information analysis

The relationship between  $PM_{10}$  concentrations and rainfall can also be determined by studying the node distribution in the MVG layers. Equation 4 shows that the joint probability between the  $PM_{10}$  and rainfall nodes  $(P(k^{[PM_{10}]}, k^{[Rainfall]}))$  is a building block of the interlayer mutual information  $(I_{PM_{10},Rainfall})$ . Thus, we first computed the joint probability before performing the interlayer mutual information analysis.

Figure 7(a) and 7(b) illustrate the quantity  $P(k^{[PM_{10}]}, k^{[Rainfall]})$  for the low dust season (October to April) and the high dust season (May to September) for the 11-y period. In these Figures, the colours indicate the probability that a node in the MVG has a degree equal to  $k^{PM_{10}}$  and  $k^{Rainfall}$ in the layers corresponding to  $PM_{10}$  and rainfall VG, respectively. Overall, the most likely combinations of k values were those below a value of 20. For

higher degrees (k > 60),  $P(k^{[PM_{10}]}, k^{[Rainfall]})$  becomes less significant. The 355 probability asymptotically approaches both the X and Y axes. According 356 to Carmona-Cabezas et al. (2020), as the degree increases, the probability of 357 finding  $k^{PM_{10}}$  and  $k^{Rainfall}$  with close values decreases exponentially. This 358 demonstrates alternation between the hubs of the two time series. Due to 359 the wet scavenging phenomenon, high daily values of  $PM_{10}$  and rainfall are 360 less likely to occur on the same day. Figure 7(a-b) shows a difference in 361 behaviour between both seasons. A concentration of probability is more 362 pronounced in the high dust season (Figure 7(b)), and higher values in a 363 low-degree area add red to the plot. In addition, the overall shape of the plot 364 shrinks for the same period and has shorter tails. These results are consistent 365 with those obtained for  $\omega$ . The impact of rainfall on  $PM_{10}$  concentrations 366 in the atmosphere more loaded with dust from May to September is more 367 efficient because of a more significant wet scavenging phenomenon (Tiwari 368 et al., 2012; Yoo et al., 2014). 369

Figure 7(c) shows the monthly  $I_{PM_{10},Rainfall}$  computed over the 11-y 370 period. The interlayer mutual information, which provides an idea of the 371 typical amount of information flow in the system (Lacasa et al., 2015), is 372 directly related to the joint probability and also measures the correlation 373 between degrees in the system. Thus, the interlayer mutual information 374 may indicate the degree of correlation among the distributions and, hence, 375 the behaviour of the two series. Because  $I_{PM_{10},Rainfall} > 1$ ,  $k^{PM_{10}}$  and 376  $k^{Rainfall}$  always have higher correlations in May ( $I_{PM_{10},Rainfall} = 1.25$ ), Au-377 gust  $(I_{PM_{10},Rainfall} = 1.14)$  and November  $(I_{PM_{10},Rainfall} = 1.15)$ . A study 378 in the Caribbean basin by Gouirand et al. (2020) showed that the averages 379 transition dates from winter to summer and from summer to winter occurred 380 on average 13 May ( $\pm$  9 days) and 26 October ( $\pm$  12 days), respectively. 381

These transition periods correspond to comparatively high  $I_{PM_{10},Rainfall}$  val-382 ues. Due to the standard deviations, the winter to summer transition always 383 occurs in May, whereas the summer to winter transition can occur in Octo-384 ber or November. This could explain why the May peak was much larger. 385 In addition, May corresponds to the beginning of the high dust season (Plo-386 coste and Pavón-Domínguez, 2020b), whereas November corresponds to the 387 end of the rainy season (Bertin and Frangi, 2013). Therefore, these periods 388 feature strong inter-layer correlations between  $k^{PM_{10}}$  and  $k^{Rainfall}$ . 389

#### 390 5. Conclusion

In conclusion, our results clearly highlight the efficiency of multilayer 391 complex networks for tracking the correlations between particulate matter 392  $(PM_{10})$  and rainfall time series. The aim of this study was to investigate the 393 wet scavenging phenomenon of  $PM_{10}$  by rainfall in the Caribbean area using 394 MVGs. We highlighted the fractal nature of both time series and found that 395 the highest degrees (hubs) are related to the highest values in the VG frame. 396 The relationship between the values and degrees of  $PM_{10}$  is less homoge-397 neous than that of rainfall due to annual intermittency. The monthly degree 398 centrality analysis indicated the seasonality of both time series. On the 1-d 399 scale, the average edge overlap ( $\omega$ ) monthly analysis highlighted the wet 400 scavenging process of  $PM_{10}$  by rainfall throughout the year. However, this 401 process seems to be more significant during the last six months of the year, 402 when the high dust and rainy seasons are juxtaposed. The joint probability 403 results between the  $PM_{10}$  and rainfall nodes according to African dust sea-404 sonality confirmed the trend observed from the  $\omega$  values. The atmosphere is 405 loaded with dust during the high dust season, and rainfall helps restore the 406

 $PM_{10}$  atmospheric balance. Thus, the overall coherence in the multivari-407 ate time series was higher from July to December. The interlayer mutual 408 information  $(I_{PM_{10}Rainfall})$  monthly analysis showed a correlation between 409  $PM_{10}$  and rainfall structures throughout the year.  $I_{PM_{10}Rainfall}$  values were 410 higher during the transition periods between winter and summer (and vice 411 versa) in the Caribbean Basin. We assume that the transition periods allow 412 the homogenisation of the multivariate time series before the usual trend 413 is resumed. To better quantify the impact of the wet scavenging process 414 on  $PM_{10}$ , a future analysis of rainwater chemistry (organic and elemental 415 carbon) related to rainfall intensity will be conducted. 416

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Figure 7: Illustration of the joint probability distribution of the degrees of both layers for (a) the low dust season (October to April) and (b) the high dust season (May to Septem-24 ber) over an 11-y period. Each isoline shows the probability that the degree is precisely  $k^{PM_{10}}$  and  $k^{Rainfall}$  at the same time node in the VG frame; (c) monthly interlayer mutual information values ( $I_{PM_{10}Rainfall}$ ) and their standard deviations computed over an 11-y period.

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### **Declaration of interests**

 $\boxtimes$  The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

□The authors declare the following financial interests/personal relationships which may be considered as potential competing interests:

## **Credit author statement**

**Thomas Plocoste:** Conceptualization, Methodology, Software, Resources, Formal analysis, Data Curation, Visualization, Validation, Investigation, Project administration, Writing - original draft, Writing – review & editing.

**Rafael Carmona-Cabezas:** Methodology, Visualization, Formal analysis, Supervision, Validation, Investigation, Writing - original draft, Writing - review & editing.

Eduardo Gutiérrez de Ravé: Supervision, Writing – review & editing.

Francisco José Jiménez-Hornero: Supervision, Writing – review & editing.

# Highlights

Multiplex visibility graphs as a useful tool for  $PM_{10}$  wet scavenging investigation

 $\mathbf{PM}_{10}$  and rainfall hubs linked to their highest values in visibility graph frame

Significant annual variation in wet deposition efficiency

Identification of winter-summer transition periods

