

1 VISIBILITY GRAPHS OF GROUND-LEVEL OZONE TIME SERIES: A MULTIFRACTAL ANALYSIS

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1 VISIBILITY GRAPHS OF GROUND-LEVEL OZONE TIME SERIES: A MULTIFRACTAL ANALYSIS

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4 ABSTRACT

5 A recent method based on the concurrence of complex networks and multifractal
6 analyses is applied for the first time to explore ground-level ozone behavior. Ozone time
7 series are converted into 2-D complex networks for their posterior analysis. The
8 searched purpose is to check the suitability of this transformation and to see whether
9 some features of these complex networks could constitute a preliminary analysis before
10 the more thorough multifractal formalism.

11 Results show effectively that the exposed transformation stores the original
12 information about the ozone dynamics and gives meaningful knowledge about the time
13 series. Based on these results, the multifractal analysis of the complex networks is
14 performed. Looking at the physical meaning of the multifractal properties (such as
15 fractal dimensions and singularity spectrum), a relationship between those and the
16 degree distribution of the complex networks is found.

17 In addition to all the promising results, this novel connection between time series
18 and complex networks can deal with both stationary and non-stationary time series,
19 overcoming one of the main limitations of multifractal analysis. Therefore, this
20 technique can be regarded as an alternative to give supplementary information within
21 the study of complex signals.

22 KEYWORDS

23 - Ground-level ozone

24 - Visibility graphs

25 - Multifractal analysis

26 - Sandbox method

27

28 1. INTRODUCTION

29 Many studies have been performed about ground-level ozone over the last decades.

30 The importance of ozone characterization and analysis lies on the fact that it is one of
31 the main photochemical oxidants (due to its abundance). This irritant gas has serious
32 repercussions for human health and harvests when its concentration is high (Doherty et
33 al., 2009). Those exposed damages have an impact from the economical point of view,
34 and according to Miao et al. (2017), they lead every year to losses of several billions of
35 dollars.

36 Ozone is a secondary pollutant whose chemical formation and destruction
37 mechanisms are known to be photochemical and nonlinear processes (Graedel and
38 Crutzen, 1993; Trainer et al, 2000). These mechanisms highly depend on meteorological
39 variables such as temperature, wind direction and mainly solar radiation (Graedel and
40 Crutzen, 1993; Guicherit and Van Dop 1977), as it has been studied for the Mediteranean
41 area in several works (Güsten et al., 1994; Kouvarakis et al., 2000; Ribas y Peñuelas
42 2004). In addition to that, ozone concentration depends as well on chemical precursors,
43 such as nitrogen oxides and volatile organic compounds derived from the urban and

44 industrial activity (Sillman, 1999). All these factors make the analysis of the temporal
45 evolution of ozone a very complex task indeed.

46 Due to the facts exposed above, ozone studies based on traditional statistical
47 analysis may provide a limited description of more complex dynamics of time series
48 where the variability is high. The reason for this limitation is that these methods
49 approximate and smooth the signal by means of fitting to functions, with the derived
50 loss of information. Besides, they base their results on one (time) scale, while the
51 physical phenomenon can appear at several scales due to the number of variables in
52 play (Zelege and Si 2006). On the contrary, multifractal methods can be used to save this
53 drawback, since they work directly with the raw data extracting the information from
54 their singularities. Furthermore, fractals (and multifractals) are characterized for being
55 self-similar when divided into smaller parts (i.e. they are scale-independent) or at least
56 their statistical properties (Mandelbrot 1982). That way, if a natural phenomenon can
57 be characterized by means of multifractal parameters, these will be able to describe it
58 for a range of scales.

59 In the presented work, a link between the multifractal analysis and complex
60 networks has been tested for the description of ground-level ozone dynamics. To that
61 purpose, ozone concentration time series have been transformed into 2-D visibility
62 graphs (VG) (whose topology inherits the features of the associated time series) and
63 then evaluated using two methods for multifractal analysis: the sandbox method in
64 order to compute the generalized fractal dimensions (Rényi spectrum) and the approach
65 introduced by Chhabra and Jensen (1989) for the calculation of the α -spectrum as an
66 independent value from the other.

67 With this study, the main purpose is to check the suitability of the multifractal
68 analysis performed over the VGs by connecting their degree distribution with Rényi and
69 α -spectra. It could be expected, since the resulting graph stores indeed much of the
70 original information and properties of the original time series (Lacasa et al., 2008).

71

72 2. MATERIALS AND METHODS

73 2.1. Data

74 The information that has been used in the analysis of this work corresponds to a 10-
75 min ozone concentration data collected for the months of January, April, July and
76 October in 2007. These time series of ozone concentration can be seen in Figure 1, for
77 the four months analyzed in this study.

78 The chosen region is the western part of Andalusia (Spain) since as exposed by
79 Domínguez-López et al. (2014), this area meets the weather conditions (high
80 temperatures and solar radiation), orographic (the valley of the Guadalquivir river) and
81 anthropic (four capitals and two important industrial centers such as the chemical pole
82 of Huelva and the Bay of Algeciras) to be potentially vulnerable to pollution by surface-
83 level ozone. The measurements were performed at the urban station located in Lepanto,
84 Córdoba (37.53° N, 4.47° W). The cited station belongs to the regional network in charge
85 of controlling the air quality in Andalusia, co-financed by the Consejería de
86 Medioambiente (Regional Environmental Department) and the European Union. This
87 station is located at 117 m of altitude and the average temperature for each month is 8,
88 15, 28 and 18 °C (January, April, July and October, respectively). The average direct solar
89 radiation is 310.2, 577.8, 991.4 and 419.3 MJ/m² (again for the four months ordered).

90 The climate of the zone of study, according to the Köppen-Geiger classification, is
91 defined as Csa with warm average temperatures and hot and dry summer.

92 As it can be seen in Figure 1, the ozone concentrations are especially high in summer
93 (July) and low in Winter (January). As it has been previously discussed for this region by
94 Adame et al. (2007) and Jiménez-Hornero et al. (2010a), that is due the fact that the
95 conditions for the ozone creation are more suitable from the end of January. The
96 progressive raise of temperature and solar radiation that reaches its peak in July allows
97 those higher creation rates and thus its concentration. One of the reactions that governs
98 the ozone production can be found below (Graedel and Crutzen, 1993).



99 It is a reversible photochemical reaction which tends to the ozone production when
100 there is energy available in form of light (right sense of the arrow) and in the other way
101 when there is not. For that reason, the higher values of ozone are always found during
102 the day and vice versa, happening the same with summer and winter respectively, as
103 discussed before.

104

105 2.2. Visibility graph

106 One of the main fields of application of the multifractal analysis is referred to the
107 study of graphs and complex networks. A graph can be defined as a set of vertices, points
108 or nodes connected to each other by lines that are usually called *edges*. A tool to
109 transform time series into a graph was presented by Lacasa et al. (2008). This new
110 complex network receives the name of *visibility graph* and has been proven to inherit

111 many of the properties of the original series (being some multifractal properties among
112 them).

113 In order to construct the visibility matrix which contains the information of all the
114 nodes of the new system, it is necessary to establish a criterion to discern whether two
115 points would be connected or not. This criterion reads as follows: two arbitrary data
116 points from the time series (t_a, y_a) and (t_b, y_b) have visibility (and would become two
117 connected nodes in the graph) if any other data point (t_c, y_c) between them ($t_a < t_c < t_b$)
118 fulfills the following condition:

$$y_c < y_a + (y_b - y_a) \frac{t_c - t_a}{t_b - t_a} \quad (2)$$

119 In Figure 2, an application of this method for a simple time series is given as
120 illustration. As it can be seen, the original time series has been transformed into a
121 complex network. The complexity of the original series is inherited by the new graph, as
122 it has been found by Lacasa et al. (2008, 2010), meaning that for instance a periodic time
123 series, would lead to a regular graph.

124 The result of applying this visibility method is a $N \times N$ adjacency binary matrix, being
125 N the number of points in the set. Each row of the matrix contains the information of a
126 different node, such as $a_{ij} = 1$ means that the node i and j have visibility; whereas
127 $a_{ij} = 0$ means that there is no edge between them. The resulting matrix has several
128 properties that can be used to simplify the algorithm and thus reduce the computational
129 required time.

- 130 • Hollow matrix: All the elements in the diagonal are zero ($a_{ii} = 0$), because there
 131 is no visibility of an element with itself, since there are no intermediate points to
 132 fulfill the criterion.
- 133 • Symmetric matrix: The elements satisfy $a_{ij} = a_{ji}$, due to the reciprocity of the
 134 visibility between two nodes.
- 135 • Nearest neighbors: Because each point always sees the closest previous and next
 136 node, the elements surrounding the diagonal are always 1 ($a_{ij} = 1$ for $j = i \pm$
 137 1).

138 Taking all of that into account, the visibility matrix A has a general form as shown
 139 below:

$$140 \quad A = \begin{pmatrix} 0 & 1 & \dots & a_{1,N} \\ 1 & 0 & 1 & \vdots \\ \vdots & 1 & \ddots & 1 \\ a_{N,1} & \dots & 1 & 0 \end{pmatrix}$$

141

142 The degree of a node (k_i) can be defined as the number nodes that have
 143 reciprocal visibility with the first one ($k_i = \sum_j a_{ij}$). In Figure 2, the degree of the first
 144 node is $k = 3$, for the second one $k = 2$, for the third one $k = 3$, and so on. From
 145 the degree of each one of the nodes present in the VG, it is possible to obtain the
 146 degree distribution of the sample ($P(k)$), which is nothing but the probability that
 147 every degree has.

148 Previous works have shown how the analysis of this degree distribution built
 149 from the VG can effectively describe the nature of the time series (Lacasa et al.,
 150 2008; Mali et al., 2018), distinguishing between periodic, random or fractal series for

151 instance. Therefore, by studying the degree distribution, one can get a first insight
 152 on the behavior of the ozone time series before stepping into a more complex
 153 multifractal analysis. As stated by Lacasa et al. (2009, 2010), time-series that have
 154 VGs with degree distributions that follow power laws such as $P(k) \propto k^{-\gamma}$, can be
 155 considered as scale free.

156

157 2.3. Multifractal measurements

158 While fractal analysis is based on the complexity of a fractal set, the multifractal
 159 approach can describe the distribution of a given measure over a fractal object
 160 (Mandelbrot, 1974; Halsey et al., 1986). It implies the possibility of having different
 161 densities depending on the region of application.

162 There are two ways of representing multifractals: the generalized fractal dimensions
 163 D_q and the singularity or multifractal spectrum ($f(\alpha)$). Both of them are discussed below
 164 separately. Typically, the multifractal analysis has been widely performed by means of
 165 the fixed-size algorithms (FSA), that rely on the subdivision of the system into smaller
 166 parts with equal size and then that size is changed iteratively. The method used in this
 167 work is the “sand-box algorithm” and will be discussed later on.

168 As stated before, the first of the measurements when it comes to multifractal
 169 analysis are the generalized fractal or Rényi dimensions D_q , which describe the scaling
 170 exponents of the q th moments of the system and can be defined (Feder, 1988) as:

$$D_q = \frac{1}{q-1} \lim_{\delta \rightarrow 0} \frac{\ln Z_q(\delta)}{\ln \delta} \quad \forall q \neq 1 \quad (3)$$

171

$$D_1 = \lim_{\delta \rightarrow 0} \frac{\sum_{i=1}^{N_c(\delta)} \mu_i(\delta) \ln \mu_i(\delta)}{\ln \delta} \quad (4)$$

172

173 With q the moment order, δ the size of the used cells to cover the system, $Z_q(\delta)$ the
 174 partition function, $N_c(\delta)$ the number of cells with length δ and $\mu_i(\delta)$ the probability
 175 measurement of each cell. The expression for D_1 is obtained taking the limit of D_q when
 176 $q \rightarrow 1$.

177 From these generalized fractal dimensions, it must be pointed out that $D_{q=0}$
 178 corresponds to the fractal dimension of the given system or *box-counting dimension*,
 179 $D_{q=1}$ is the so-called *information entropy* and $D_{q=2}$ the *correlation dimension*. The limits
 180 of D_q when q goes to $-\infty$ and $+\infty$ describe the scaling properties of the regions where
 181 the measure is more rarified and concentrated, respectively.

182 The other set of multifractal parameters is the so-called singularity spectrum ($f(\alpha)$),
 183 as commented previously. A frequent method to determine is based on the use of a
 184 Legendre transform from mass exponents $\tau(q)$. However, some authors as Chhabra and
 185 Jensen (1989) and Veneziano et al (1995), state that it can lead to some errors due to
 186 the inclusion of spurious points and error amplification from the derivative.
 187 Furthermore, it does not yield independent measurements from the Rényi spectrum, as
 188 $\tau(q) = (1 - q)D_q$. As an alternative to this and more focused on experimental data,
 189 Chhabra and Jensen (1989) proposed a direct method to determine the α -spectrum,
 190 overcoming the drawbacks referred before. This technique relies on the normalized
 191 measure $\beta_i(q)$, μ_i in the original work, for computing the probabilities of the boxes of
 192 radius r :

$$\beta_i(q, r) = [P_i(r)]^q / \sum_j [P_j(r)]^q \quad (5)$$

193 With $P_i(r)$ the different fractal measurements for each box of radius r (number of
194 nodes in this case). From this, $f(\alpha)$ and α can be retrieved using the next expressions:

$$f(q) = \lim_{r \rightarrow 0} \frac{\sum_i \beta_i(q, r) \log[\beta_i(q, r)]}{\log r} \quad (6)$$

195

$$\alpha(q) = \lim_{r \rightarrow 0} \frac{\sum_i \beta_i(q, r) \log[P_i(r)]}{\log r} \quad (7)$$

196 Being α known as the Lipschitz-Hölder exponent. In practice, those quantities are
197 computed from the slope of $\sum_i \beta_i(q, r) \log[\beta_i(q, r)]$ over $\log r$ for $f(q)$; and
198 $\sum_i \beta_i(q, r) \log[P_i(r)]$ over $\log r$ for $\alpha(q)$. This slope is determined by means of a linear
199 regression in the same range of radii where the other fractal measures are computed.

200

201 2.4. Sandbox algorithm

202 The sandbox algorithm (SBA) was firstly introduced by Tél et al. (1989) and
203 developed by Vicsek et al. (1990), as an improvement of the previously used fixed-size
204 box-counting methods for computing the generalized fractal dimensions. The main
205 advantage of this method with respect to other box-counting FSA is that it is capable of
206 properly determine the side corresponding to the negative values of q from both the
207 Rényi and singularity spectra.

208 The basic idea behind the SBA is that for each radius, a number of randomly placed
209 boxes are selected, and they are always centered in a non-zero point of the system (a

210 node). In that way the entire network is covered with those boxes by choosing a
 211 sufficiently high number of them. For each box (B), the probability measurement used
 212 is computed as shown in equation (8).

$$\mu(B) = \frac{M(B)}{M_0} \quad (8)$$

213 Once that quantity is computed for each sandbox for a given radius, the generalized
 214 fractal dimensions can be obtained as explained previously in the multifractal
 215 measurement part. Applying it to equation (3), the following formula is obtained for
 216 D_q^{sb} :

$$D_q^{sb} = \frac{1}{q-1} \lim_{r \rightarrow 0} \frac{\ln \langle \mu(B)^{q-1} \rangle}{\ln r} \quad \forall q \neq 1 \quad (9)$$

217 And the expression when $q = 1$ for the SBA also can be adapted from equation (4):

$$D_1^{sb} = \lim_{r \rightarrow 0} \frac{\langle \ln \mu(B) \rangle}{\ln r} \quad (10)$$

218 With the aim of implementing the SBA on the VG, the steps to follow are:

- 219 1) The original time series is transformed into a VG, resulting on a matrix as
 220 described in the previous section.
- 221 2) A range of different radii is set in order to cover the entire network. These
 222 radii are chosen between 1 and N, being N the total number of nodes. For
 223 this case, the typical values are $r \in [1, 100]$, because larger values have
 224 proven to give the same information, since the curves start to saturate.
- 225 3) Then, N_c centers are selected for each radius. This number is inversely
 226 proportional to the radius itself, because the greater the box is, the smaller

227 number of boxes will be needed to cover the network. The location of the
228 centers is randomly chosen within the nodes of the graph.

229 4) For each sandbox, the amount of nodes inside the box connected to the
230 center (M) are counted, giving N_c values of $M(B)$. From here, the quantity
231 $\mu(B)$ (for every box) is computed by means of dividing the previous $M(B)$ by
232 the total amount of nodes connected to the center M_0 . Then $\mu(B)$ is used to
233 compute the partition function for all the q values chosen, being the average
234 in equations (9) and (10) over all the sandboxes created.

235 5) The steps 3) and 4) are then repeated for all the radii considered, obtaining
236 a value of the partition function for each one of them and q .

237

238 3. RESULTS AND DISCUSSION

239 3.1. Degree distribution

240 Firstly, a fast method to check the fractal nature of the ozone concentration time
241 series, before performing a deeper study, is described here. This method consists on
242 analyzing the degree distribution of the VGs.

243 The mentioned distribution is computed as the number of nodes that have a
244 given degree and divided by the total number of them in the VG. It is clear that the
245 greater the degree is, the less likely to be repeated within the network it will be;
246 because large degrees are exclusive of nodes with the highest concentration (which
247 we will refer as *hub*) due to their typically high visibility. As those hubs correspond
248 normally to the extreme values of ozone concentration of each day, their likeliness
249 will be lower, since the most repeated ones will be the ones close to the average.

250 That leads to an expected distribution with a negative trend, as can be seen in Figure
251 3.

252 Once the degree distributions of the different months are computed, a clear
253 fractal behavior is observed since all of them follow a power law $P(k) \propto k^{-\gamma}$ as
254 expected, with a linear part in the last part of the log-log plot. After computing the
255 slope for the values of $k \geq 30$ in all the cases, the γ parameter can be determined,
256 obtaining that the biggest and the lowest ones correspond to July and January
257 respectively. Although this is a fast and direct method to determine whether the
258 series is fractal or not, it does not give much detailed information about the
259 differences between each month, since all of them share a very similar degree
260 distribution. Hence a deeper analysis devoted to the multifractal properties of the
261 series has been done with that aim, as will be presented in the following section, in
262 order to give some light to the usability of this degree distribution and some of its
263 parameters.

264

265 3.2. Multifractality

266 In this part, the authors present two independent methods to study the
267 multifractality of the signal, both previously described: the generalized fractal
268 (Rényi) dimensions and the α spectra.

269 For the analysis of the Rényi dimensions of the samples, the quantities
270 $\ln\langle\mu(B)^{q-1}\rangle / (q - 1)$ for $q \neq 1$ and $\langle\ln \mu(B)\rangle$ for $q = 1$ are plotted against $\ln r$, as
271 seen in Figure 4. The interval used for the radii goes from $r = 1$ to $r = 100$, as it was
272 proven to be enough in order to reproduce the expected linear behavior (after

273 several tests). It must be pointed out that the interval taken for the linear regression
274 is always selected so that the Pearson correlation coefficient is $r \geq 0.99$. For April,
275 July and October, the chosen range has been from $\ln r = 3$ up to end of the curve;
276 whereas for January the range had to be taken such as $\ln r \in [2, 4]$ because the
277 linear part is larger and to avoid an artifact that appears for higher radii and made
278 the results misleading (can be observed in Figure 7). In all those cases, the values
279 correspond with the higher radii used in the SBA, in accordance with previous studies
280 that used the same technique (Ariza-Villaverde et al., 2015; De Bartolo et al., 2004;
281 Jiménez-Hornero et al. 2013). After the linear regressions are performed for all
282 values of q of each month, the Rényi spectrum can be constructed following
283 equations (9) and (10). In Figure 5 the result of such procedure is shown in the form
284 of the generalized fractal dimensions. After observing the graph, it can be inferred
285 that the VG properly reproduces the multifractal behavior of the series, which was
286 previously demonstrated to be so (He, 2017; Jiménez-Hornero et al., 2010a, 2010b;
287 Pavón-Domínguez et al., 2013), since for all the studied months $D_0 > D_1 > D_2$ (see
288 Table 1). The difference between the maximum value of D_q and the minimum (ΔD_q)
289 is usually taken as a measure of the multifractality degree of the signal (ozone
290 concentration in this case) (Ariza-Villaverde, 2013; Telesca et al., 2004). Looking at
291 Table 1, this degree seems to be sensibly higher for July than the rest of the months,
292 followed by April and October (with a very similar value) and then the last one is
293 January.

294 Now, the fractal dimensions themselves from the Rényi spectrum (D_0 , D_1 and
295 D_2), are being discussed. Firstly, the capacity dimension or “box-counting”
296 dimension D_0 is related with the number of boxes needed in order to cover the

297 fractal object. The more boxes are needed, the more extended is the fractal object
298 and then the greater would be D_0 . Since in this case the fractals are different for each
299 iteration (each node sees a different visibility network), the final behavior would be
300 expected to be an average of the whole system. In the context of these complex
301 networks, a high extension of the object would mean a bigger degree (then related
302 to hubs). Following that logic, a higher value of D_0 is related with greater number of
303 hubs in the system, meaning that one could expect the average degree \bar{k} of the
304 distribution increased. This is found in the results as can be observed in Table 1.
305 Values for the months of January and April are very close to each other, while the
306 one for July is sensibly the highest and October has an intermediate value. This result
307 is in accordance to the behavior of the ozone concentration, since for instance the
308 month of July is the one with the greatest temperature and UV radiation rates and
309 therefore elevated values are reached more often, producing a major number of
310 hubs in the VG.

311 The second parameter is the entropy dimension D_1 , which is usually known as
312 the dimension that is related to the uniformity of the data and how different is the
313 probability of certain regions to be visited by a randomly chosen box with respect to
314 others. A way of measuring its uniformity is looking at the difference between D_0
315 and D_1 , because when they are equal it means that the sample is uniform. The
316 greater ($D_0 - D_1$), the less uniform it is. Also, another reason to take the difference
317 instead of the absolute value is because we are interested on comparing the curves
318 and their behavior, and since the D_0 of each one is different, it is necessary to
319 establish a reference point. When it is translated to the context of VG, a greater
320 uniformity would mean less difference between the degrees of the sample, hence

321 the standard deviation of the degree would be decreased. In that case, what should
322 be observed in the original ozone data is that the differences encountered in the
323 data are less significant (what is indeed observed for the months with lower $(D_0 -$
324 $D_1)$). The most uniform would be January and again the other extreme is found for
325 July. The reason for this is that in July the differences encountered between day and
326 night concentrations of ozone are much greater than the other months, since during
327 the night they remain at their minimum values for all months. In contrast, the
328 maximum values are much higher in summer than the rest of the year because of
329 the higher UV exposures and the opposite case for winter.

330 The last parameter that has been computed from the Rényi spectrum is the
331 correlation dimension D_2 . For the same objective reason as in $(D_0 - D_1)$, here $(D_0 -$
332 $D_2)$ values are discussed as well. Once again, January and July exhibit the extreme
333 cases whereas the other two months are located in between. In this case the authors
334 have not been able to directly relate this magnitude with any property of the degree
335 distribution of the VGs as for $(D_0 - D_1)$ and D_0 .

336 As commented before, another important feature that can be studied is the
337 spectrum of α , following the equations (6) and (7). Applying a similar approach to
338 the one used to compute D_q , a linear regression is performed for each value of q to
339 compute the corresponding magnitude (α in Figure 6 and $f(\alpha)$ in Figure 7). For each
340 month, the range for the regression has been chosen to be the same that was used
341 for the Rényi dimensions, in order to conserve the scale used. In this case, the
342 obtained curves have as well a positive trend that increases with q for the case of
343 the curves of Figure 6; and first increases and then decreases for Figure 7 (as

344 expected, because $f(\alpha)$ should have a maximum and then drop). As commented
345 before, the artifact that made us change the chosen interval for the regression in
346 January can be observed in Figure 7 for high values of radius.

347 Once both α and $f(\alpha)$ are resolved for all the possible values of q , the spectrum
348 of every month can be plotted, being the ones shown in Figure 8. All the spectra as
349 can be seen have their maximum at $f(\alpha) = 1$ and start at $\alpha = 0$, while the overall
350 shape depends on each case. Several properties of the underlying signal can be
351 extracted from the spectra. The width of the curve W is shown in Table 1, as well as
352 the position of the curve maximum α_0 . As in the case of generalized fractal
353 dimensions D_q , July is the month with the widest spectrum and highest α_0 . January
354 as well exhibits the lowest values, whereas April and October have intermediate
355 ones. This width W is related with the multifractality degree of the signal (Ariza-
356 Villaverde, 2013; Telesca et al., 2004).

357 According to the shapes of the curves, the spectra are strongly non-symmetric,
358 being the right tail much more pronounced than the left one for all the different
359 months. The right side is usually associated to the homogeneity of low values in a
360 temporal distribution of data, whereas the left one is related with the same feature
361 of the high values instead. Therefore, the behavior of the four months is very similar
362 for high values (left tail) while it differs significantly for low ones (right tail). In Figure
363 8, all of them show an heterogenous behavior since the α distribution is not uniform
364 along the curve, being concentrated on the left and right extremes. Nevertheless, it
365 is possible to extract from that figure that this heterogeneity is more pronounced in
366 July. This fact suggests an influence of the higher UV radiation that creates a greater

367 difference between day and night ozone concentrations. This is in accordance with
368 what was shown in the previous analysis described above for $(D_0 - D_1)$.

369

370 4. CONCLUSIONS

371 The results of this work show that the multifractal analysis of VGs from ground-
372 level ozone concentration time series is a suitable tool to describe the seasonal
373 dynamics of this air pollutant. VGs have proven to have several advantages such as:
374 i) their topology inherits the features of the associated time series, which ends up
375 resulting on supplementary information through the degree distribution; ii) they can
376 be use for both stationary and non-stationary time series, removing the multifractal
377 analysis requirement of dealing only with stationary processes. Besides, VGs filter
378 out trends in the signal, making unnecessary to apply detrended methods. iii) Also,
379 they can be applied to multivariate time series, which can be very helpful in order to
380 find correlations between tropospheric ozone and its precursors for instance; iv) and
381 finally, this novel connection between time series and complex networks opens a
382 broad range of possibilities within the study of complex signals.

383 When it comes to the multifractal analysis performed in this work, both the SBA
384 and the Chhabra and Jensen method for the Rényi and singularity spectra
385 respectively, where chosen based on their advantages with respect to other. The
386 SBA overcomes the drawbacks of the box-counting algorithm for the computation
387 of the generalized fractal dimensions for negative probability moment orders q .
388 Furthermore, the Chhabra and Jensen method for the α -spectrum does not need a

389 Legendre transform to be applied (with the associated error to it) and gives and
390 independent measurement from the SBA as well, being more robust for comparison.

391 After applying the methods mentioned above, clear and coherent results that fits
392 the expected behavior of the ozone dynamics were found for the different months.
393 Furthermore, several properties can be directly inferred from the degree
394 distribution almost at first sight, meaning a powerful tool for predicting results
395 before any more complex data treatment is performed. The same relations between
396 multifractal parameters such as capacity dimension, W of the α -spectrum amongst
397 others are found in simple statistical parameters as the average or standard
398 deviation of the degree distribution of the VG.

399 To conclude, authors would like to point out that this promising technique could
400 be extended to other applications due to the many possibilities that complex
401 networks have. One possible target would be the relation between ozone and its
402 precursors, which could be looked by means of relating the VG of each one
403 separately or using multi-layer networks as proposed by Lacasa et al. (2015).

404

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410

411 6. REFERENCES

412 Adame, J.A., Lozano, A., Bolívar, J.P., De la Morena, B.A., Contreras, J., Godoy, F., 2008.

413 Behavior, distribution and variability of surface ozone at an arid region in the
414 south of Iberian Peninsula (Seville, Spain). *Chemosphere* 70, 841–849.
415 <https://doi.org/10.1016/j.chemosphere.2007.07.009>

416

417 Ariza-Villaverde, A.B., Jiménez-Hornero, F.J., Gutiérrez de Ravé, E., 2015. Influence of
418 DEM resolution on drainage network extraction: A multifractal analysis.
419 *Geomorphology* 241, 243–254.
420 <https://doi.org/10.1016/j.geomorph.2015.03.040>

421

422 Ariza-Villaverde, A.B., Jiménez-Hornero, F.J., Gutiérrez de Ravé, E., 2013. Multifractal
423 analysis applied to the study of the accuracy of DEM-based stream derivation.
424 *Geomorphology* 197, 85–95.
425 <https://doi.org/10.1016/j.geomorph.2013.04.040>

426

427 Chhabra, A., Jensen, R.V., 1989. Direct determination of the $f(\alpha)$ singularity spectrum.
428 *Physical Review Letters* 62, 1327–1330.
429 <https://doi.org/10.1103/PhysRevLett.62.1327>

430

431 De Bartolo, S.G., Gaudio, R., Gabriele, S., 2004. Multifractal analysis of river networks:
432 Sandbox approach. *Water Resources Research* 40.
433 <https://doi.org/10.1029/2003WR002760>

434

435 Feder, J., 1988. *Fractals, Physics of Solids and Liquids*. Springer US.

436

- 437 Graedel, T.E., Crutzen, P.J., 1993. Atmospheric change: an earth system perspective.
438 Atmospheric change: an earth system perspective.
439
- 440 Guicherit, R., van Dop, H., 1977. Photochemical production of ozone in Western Europe
441 (1971–1975) and its relation to meteorology. *Atmospheric Environment* (1967)
442 11, 145–155.
443 [https://doi.org/10.1016/0004-6981\(77\)90219-0](https://doi.org/10.1016/0004-6981(77)90219-0)
444
- 445 Güsten, H., Heinrich, G., Weppner, J., Abdel-Aal, M.M., Abdel-Hay, F.A., Ramadan, A.B.,
446 Tawfik, F.S., Ahmed, D.M., Hassan, G.K.Y., Cvitaš, T., Jeftić, J., Klasinc, L., 1994.
447 Ozone formation in the greater Cairo area. *Science of The Total Environment*
448 155, 285–295.
449 [https://doi.org/10.1016/0048-9697\(94\)90507-X](https://doi.org/10.1016/0048-9697(94)90507-X)
450
- 451 Halsey, T.C., Jensen, M.H., Kadanoff, L.P., Procaccia, I., Shraiman, B.I., 1986. Fractal
452 measures and their singularities: The characterization of strange sets. *Phys. Rev.*
453 *A* 33, 1141–1151.
454 <https://doi.org/10.1103/PhysRevA.33.1141>
455
- 456 He, H., 2017. Multifractal analysis of interactive patterns between meteorological
457 factors and pollutants in urban and rural areas. *Atmospheric Environment* 149,
458 47–54. <https://doi.org/10.1016/j.atmosenv.2016.11.004>
459
- 460 Jiménez-Hornero, F.J., Ariza-Villaverde, A.B., De Ravé, E.G., 2013. Multifractal

461 description of simulated flow velocity in idealised porous media by using the
462 sandbox method. *Fractals* 21, 1350006.
463 <https://doi.org/10.1142/S0218348X13500060>

464

465 Jiménez-Hornero, F. J., Gutiérrez de Ravé, E., Ariza-Villarverde, A.B., Giráldez, J.V.,
466 2010a. Description of the seasonal pattern in ozone concentration time series by
467 using the strange attractor multifractal formalism. *Environmental Monitoring*
468 *and Assessment* 160, 229–236.
469 <https://doi.org/10.1007/s10661-008-0690-y>

470

471 Jiménez-Hornero, Francisco J., Jiménez-Hornero, J.E., Gutiérrez de Ravé, E., Pavón-
472 Domínguez, P., 2010b. Exploring the relationship between nitrogen dioxide and
473 ground-level ozone by applying the joint multifractal analysis. *Environmental*
474 *Monitoring and Assessment* 167, 675–684.
475 <https://doi.org/10.1007/s10661-009-1083-6>

476

477 Kouvarakis, G., Tsigaridis, K., Kanakidou, M., Mihalopoulos, N., 2000. Temporal
478 variations of surface regional background ozone over Crete Island in the
479 southeast Mediterranean. *Journal of Geophysical Research: Atmospheres* 105,
480 4399–4407. <https://doi.org/10.1029/1999JD900984>

481

482 Lacasa, L., Luque, B., Ballesteros, F., Luque, J., Nuño, J.C., 2008. From time series to
483 complex networks: The visibility graph. *Proceedings of the National Academy of*
484 *Sciences* 105, 4972–4975.

485 <https://doi.org/10.1073/pnas.0709247105>

486

487 Lacasa, L., Luque, B., Luque, J., Nuño, J.C., 2009. The visibility graph: A new method for
488 estimating the Hurst exponent of fractional Brownian motion. EPL (Europhysics
489 Letters) 86, 30001.

490 <https://doi.org/10.1209/0295-5075/86/30001>

491

492 Lacasa, L., Toral, R., 2010. Description of stochastic and chaotic series using visibility
493 graphs. Physical Review E 82.

494 <https://doi.org/10.1103/PhysRevE.82.036120>

495

496 Mali, P., Manna, S.K., Mukhopadhyay, A., Haldar, P.K., Singh, G., 2018. Multifractal
497 analysis of multiparticle emission data in the framework of visibility graph and
498 sandbox algorithm. Physica A: Statistical Mechanics and its Applications 493,
499 253–266.

500 <https://doi.org/10.1016/j.physa.2017.10.015>

501

502 Mandelbrot, B.B., 1982. The fractal geometry of nature. W.H. Freeman, San Francisco.

503

504 Pavon-Dominguez, P., Jiménez-Hornero, F.J., Gutierrez de Rave, E., 2013. Multifractal
505 analysis of ground-level ozone concentrations at urban, suburban and rural
506 background monitoring sites in Southwestern Iberian Peninsula. Atmospheric
507 Pollution Research 4, 229–237.

508 <https://doi.org/10.5094/APR.2013.024>

509

510 Ribas, À., Peñuelas, J., 2004. Temporal patterns of surface ozone levels in different
511 habitats of the North Western Mediterranean basin. *Atmospheric Environment*
512 38, 985–992.

513 <https://doi.org/10.1016/j.atmosenv.2003.10.045>

514

515 Sillman, S., 1999. The relation between ozone, NO_x and hydrocarbons in urban and
516 polluted rural environments. *Atmospheric Environment* 33, 1821–1845.

517 [https://doi.org/10.1016/S1352-2310\(98\)00345-8](https://doi.org/10.1016/S1352-2310(98)00345-8)

518

519 Tél, T., Fülöp, Á., Vicsek, T., 1989. Determination of fractal dimensions for geometrical
520 multifractals. *Physica A: Statistical Mechanics and its Applications* 159, 155–166.

521 [https://doi.org/10.1016/0378-4371\(89\)90563-3](https://doi.org/10.1016/0378-4371(89)90563-3)

522

523 Telesca, L., Colangelo, G., Lapenna, V., Macchiato, M., 2004. Fluctuation dynamics in
524 geoelectrical data: an investigation by using multifractal detrended fluctuation

525 analysis. *Physics Letters A* 332, 398–404.

526 <https://doi.org/10.1016/j.physleta.2004.10.011>

527

528 Trainer, M., Parrish, D.D., Goldan, P.D., Roberts, J., Fehsenfeld, F.C., 2000. Review of
529 observation-based analysis of the regional factors influencing ozone
530 concentrations. *Atmospheric Environment* 17.

531

532 Veneziano, D., Moglen, G.E., Bras, R.L., 1995. Multifractal analysis: Pitfalls of standard

533 procedures and alternatives. Phys. Rev. E 52, 1387–1398.

534 <https://doi.org/10.1103/PhysRevE.52.1387>

535

536 Vicsek, T., Family, F., Meakin, P., 1990. Multifractal Geometry of Diffusion-Limited

537 Aggregates. Europhysics Letters (EPL) 12, 217–222.

538 <https://doi.org/10.1209/0295-5075/12/3/005>

539

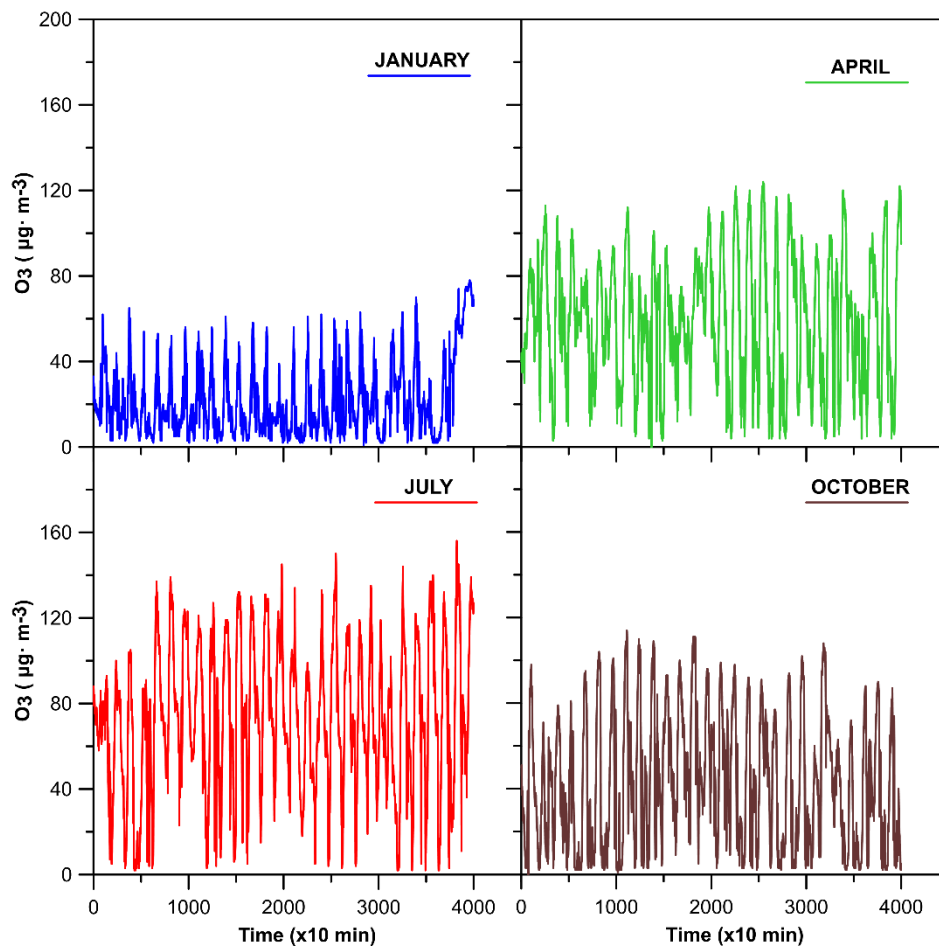
540 Zeleke, T.B., Si, B.C., 2006. Characterizing scale-dependent spatial relationships between

541 soil properties using multifractal techniques. Geoderma 134, 440–452.

542 <https://doi.org/10.1016/j.geoderma.2006.03.013>

543

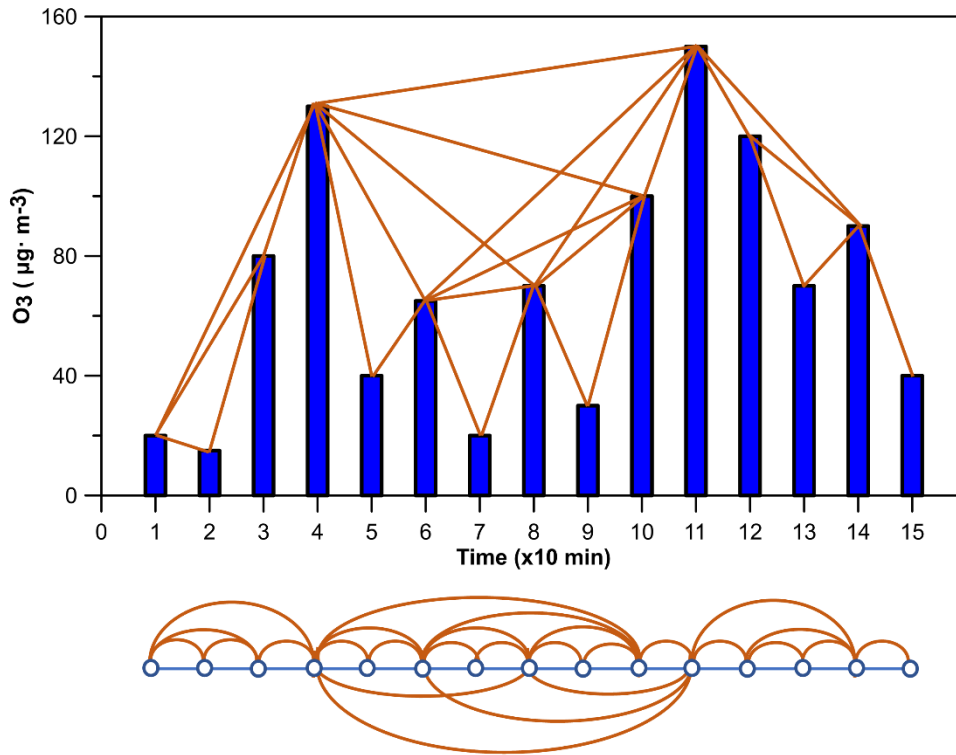
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545 7. FIGURES

546 8.

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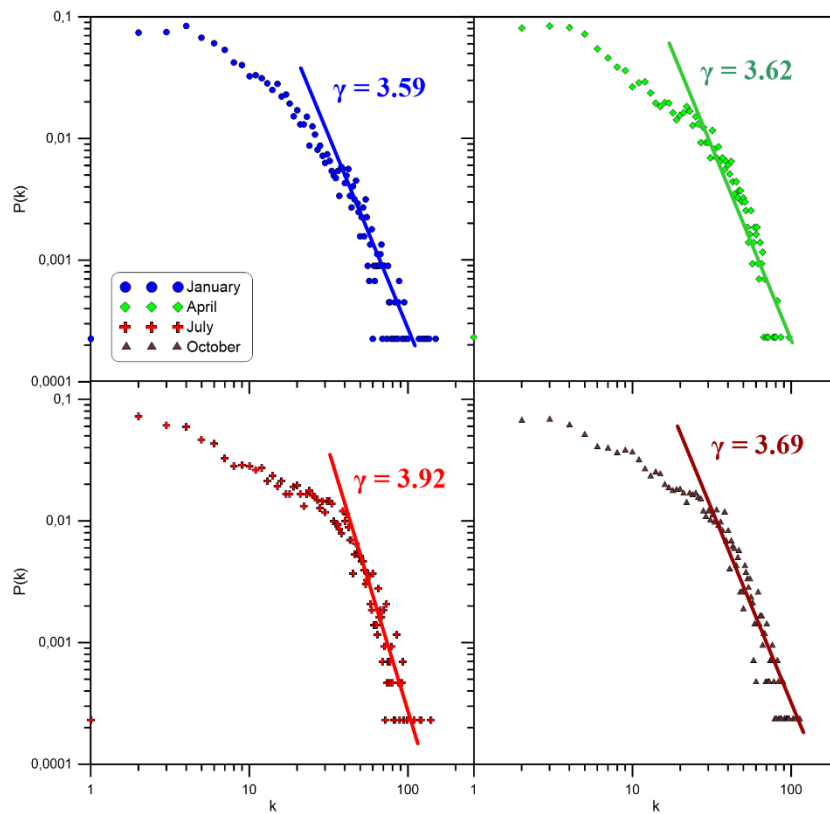
Figure 1: Ozone time series for the different analyzed months.



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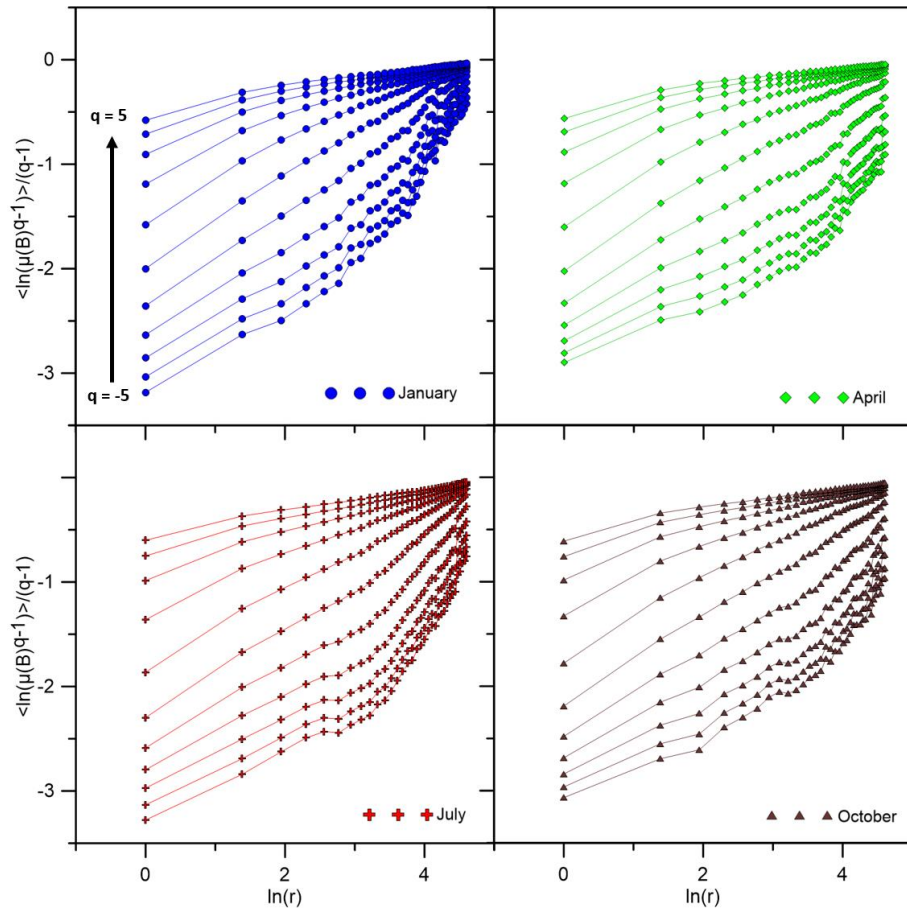
549 *Figure 2: Sample time series transformed into a complex network through the visibility*
 550 *algorithm. Below, all the connections are shown in a more visual way.*

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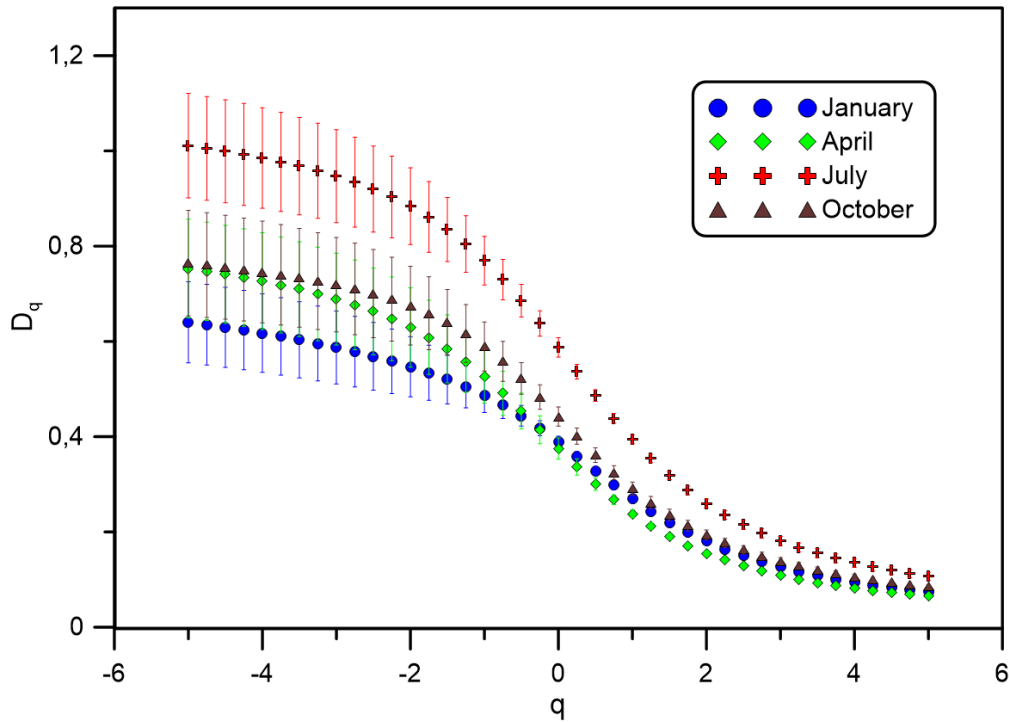
553 *Figure 3: Degree distribution of the visibility graph from each month in logarithmic*
 554 *scale. As it can be observed, the tail of the distribution shows a scale-free behavior*
 555 *because it can be fitted by a power law.*



556

557 *Figure 4: Curves of the partition function against $\ln(r)$ obtained after applying the*
 558 *sandbox algorithm.*

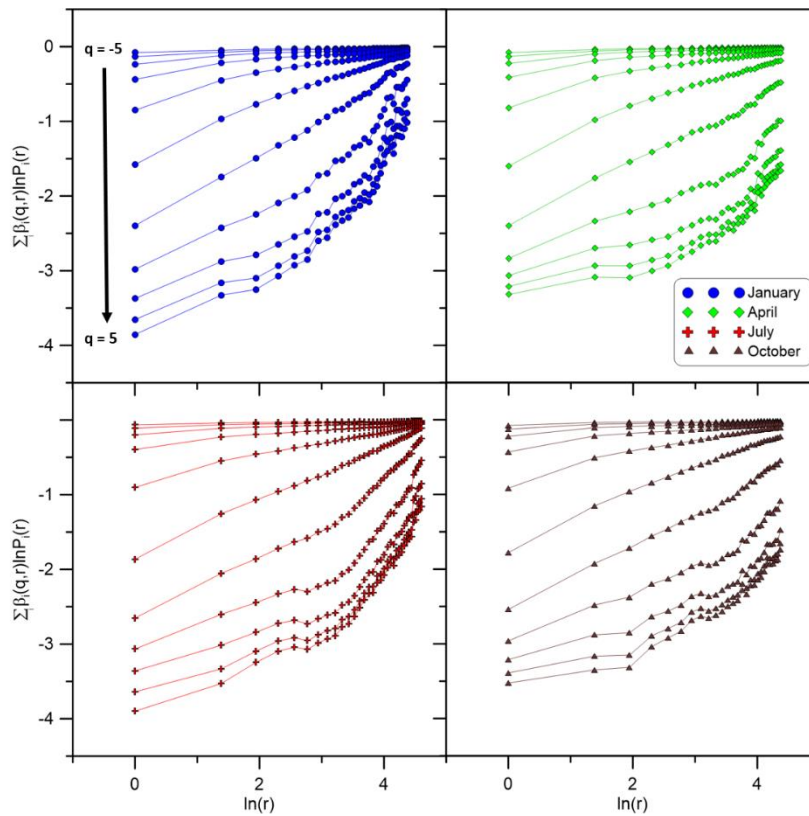
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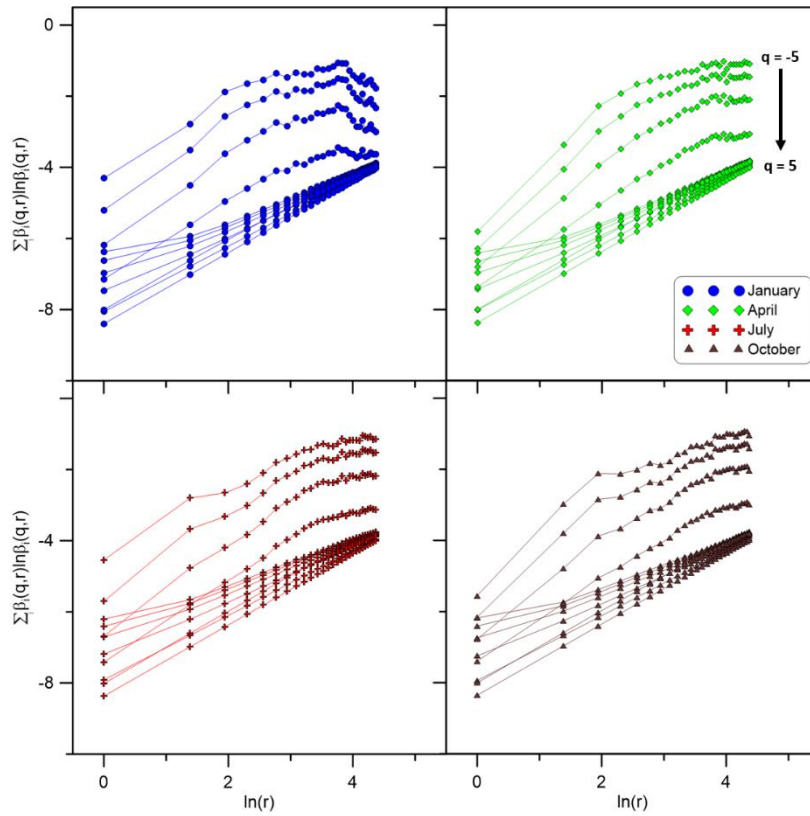
Figure 5: Rényi dimensions for each month.



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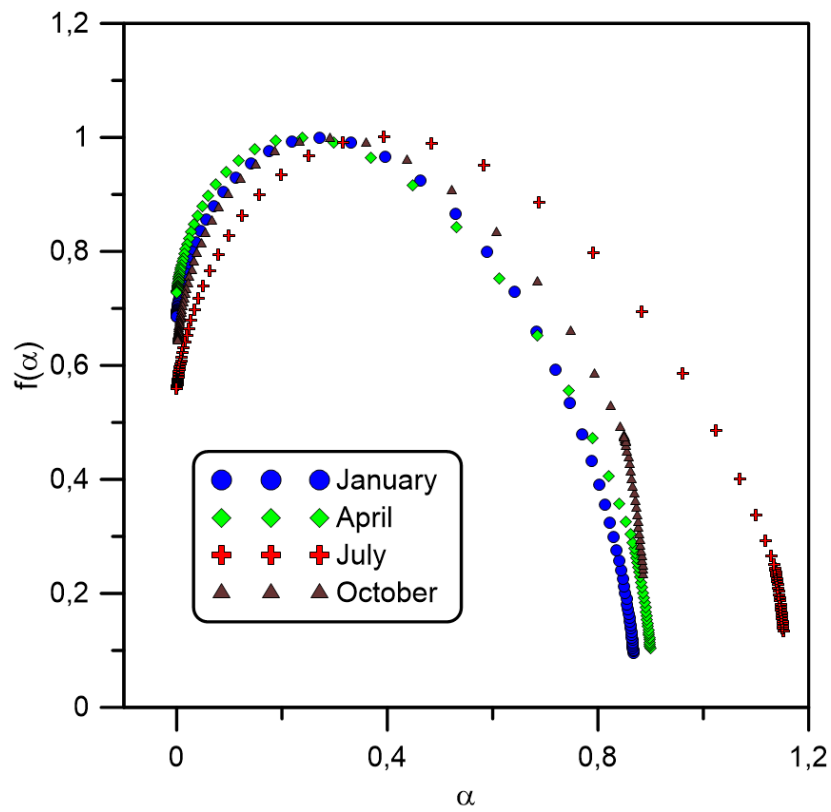
Figure 6: Curves from where α is computed by regression.



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Figure 7: Curves from where $f(\alpha)$ is computed using linear regressions.



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Figure 8: α -spectra of all the months.

569 9. TABLES

570

Table 1: Relevant values for each month.

MONTH	Average Direct Radiation (MJ/m ²)	γ	\bar{k}	σ_k	D_0	$D_0 - D_1$	$D_0 - D_2$	ΔD_q	α_0	W
January	310.2	3.59	14.78	15.00	0.389±0.012	0.12±0.03	0.21±0.02	0.56±0.10	0.291	0.779
April	577.8	3.62	15.12	13.98	0.37±0.03	0.13±0.04	0.22±0.04	0.69±0.12	0.246	0.934
July	991.4	3.92	19.94	17.69	0.59±0.02	0.20±0.04	0.33±0.03	0.90±0.12	0.443	1.152
October	419.3	3.69	17.90	15.91	0.44±0.02	0.15±0.04	0.25±0.04	0.68±0.12	0.333	0.841

571

1 **HIGHLIGHTS**

- 2 - Ground-level ozone concentration time series have a multifractal behavior.
- 3 - Visibility graphs can be used to analyze multifractality of ozone time series.
- 4 - Many aspects of ozone dynamics can be observed through the degree
- 5 distribution.
- 6 - This technique gives supplementary information within the study of complex
- 7 signals.

